The influence of temperature—moisture coupling on the occurrence of compound dry and hot events over South America: historical and future perspectives

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1 Introduction



Fig 1. Geographical scope of South America (SA) with it's climatic sub-regions defined by [3] and the moist and non moist ecoregions defined by [4].

Changes in the relationship between temperature and moisture in S. America (SA) are foreseen in a climate change scenario due to a temperature reshaping of and distributions. This precipitation resulting from association the the interaction between (i) contribution (AC), atmospheric evaluated the through here correlation between temperature and precipitation [1], and the (ii) land-atmospheric contribution (LAC), conceptualized as the soil moisture-temperature coupling [2], modulates the occurrence of and hot (CDH) compound dry extremes.

4 CDH conditions in South America

• All regions are expected to witness an increase in the mean number of summer days under CDH conditions, with particular emphasis to NWS and NSA.



Fig. 3 a and b Boxplots depicting the statistical distribution of the mean summer CDH days per year over the SA sub-regions and within the respective Wet (a) and non Wet ecoregions (b). c Percentage of increase in the mean summer CDH days per year for RCP2.6 and RCP8.5 future scenarios in respect to the Historical period. d and e Spatial distribution over SA of the variation in the mean summer CDH days per year for the RCP2.6 and RCP8.5 in respect to the historic.

2 Data and Methods • Data ERA5 reanalysis Historic (1970-2005) RCP2.6 + RCP8.5 (2006-2099) • CDH days **Dry Month** (SPI < -1)Heatwave day (CTX90pct)

Multivariate Regression Analysis

 $Y = \beta_0 + \beta_1 X_1 + \beta_2 X_2 + \varepsilon$ *Y* = Mean summer CDH days per year $\beta_1 X_1 = \beta_{AC} AC \mid \beta_2 X_2 = \beta_{LAC} LAC$

5 Influence of LAC and AC on CDH days

- In general, the level of influence of soil moisture-temperature coupling ($\beta(LAC) > 0$) on summer CDH conditions is expected to increase in future for all regions.
- A simultaneous strengthening of the influence of soil moisturetemperature coupling ($\beta(LAC) > 0$) and of compound hot and low precipitation conditions ($\beta(AC) < 0$) is only observed over NWS.



Fig. 4 Regression coeficients ($\beta \times 10^{-1}$) obtained from the multivariate regression models applied for the SA sub-regions and within the respective Wet (left panel) and non Wet (right panel) ecoregions for the historical, RCP2.6 and RCP8.5 experiments.

References

- Berg, A. et al. Interannual Coupling between Summertime Surface Temperature and Precipitation over Land: Processes and Implications for Climate Change. J. Clim. 28, 1308–1328 (2015). Gevaert, A.I. et al. Soil Moisture-Temperature Coupling in a Set of Land Surface Models. J.
- Geophys. Res. Atmos. 123(3), 1481–1498 (2018).
- Castellanos, E. et al. Central and South America. In: Climate Change 2022: Impacts, Adaptation and Vulnerability. Contribution of Working Group II to the Sixth Assessment Report of the Intergovernmental Panel on Climate Change. (2022)
- Olson, D.M. et al. Terrestrial Ecoregions of the World: A New Map of Life on Earth: A new global map of terrestrial ecoregions provides an innovative tool for conserving biodiversity. *BioScience*, 51, 933-938 (2001).

CORDEX (SAM22) multi-variable ensemble Historic (1970-2005)

> **Atmospheric Contribution (AC)** $\rho(P, T_{proxy})$

Land-Atmospheric Contribution (LAC) $\Pi = \rho(Rn - \lambda E, T) - \rho(Rn - \lambda Ep, T)$

NWS	Non NSA	WET NES	SAM	
				HIST
				RCP2.6
				RCP8.5

3 AC an	dl	A.
	0.8	
	0.6	۱ • ۲ ©
Future changes in the	0.4	O F
bivariate distribution	U 0.2	
of AC and LAC levels	0	
are region dependent:	-0.2	
	-0.4	
1 ΔC strengthening.	0.8	
$\mathbf{L} = \frac{\mathbf{A} \mathbf{C} \mathbf{S} \mathbf{U} \mathbf{C} \mathbf{C} \mathbf{C} \mathbf{C} \mathbf{C} \mathbf{C} \mathbf{C} C$	0.6	
	0.4	
NES WEI	C 10.2	
 NES non WET 	0	
	-0.2	
2. LAC strengthening :	-0.4	
NWS WET	0.4	
NES WET	0.2	
 NES non \/FT 	0.2	
	7	
• SAIVI VVEI	0	
 SAM no WET 	0.0	
	0.2	
3. (1) + (2):	0.6	
 NES WET 	0.4	
 NES non WET 	0.2 FAC	
	0	
	-0.2	

Fig. 2 Bivariate Kernel distributions of LAC and AC levels across different SA subregion and within the respective Wet and non Wet ecoregions for the historical period (green shades), the RCP2.6 (blue contours) and the RCP8.5 (orange contours) experiments.

Summary

- LAC levels over SA.

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• A high sptatial variability is expected in future changes of the AC and

The expected increase of CDH conditions is partially explained by the sum of two factors: (1) increase in the AC and LAC levels and (2) changes in the influence of AC and LAC on the ocurrence of CDH events.







